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flora suggest that late-glacial forest refugia were more developed here at middle altitude, rather than at higher altitude as previously suggested. The forest developed after 9800 yr B.P., while the water level remained high in the Korçë basin until 5000 yr B.P. Different environmental conditions, characterized by lower available moisture and warmer winters, progressively took place after this date. Human activity in the Korçë basin ca. 4500 yr B.P. was coeval with conditions characterized by an increase in winter temperatures and a decrease in summer moisture. © 2000 University of Washington.

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The post-glacial history of two adjacent sites in the Harvard Forest, a 10-ha swamp (Black Gum Swamp) and a 0.006-ha hollow (Hemlock Hollow) in a Tsuga canadensis forest were investigated using pollen analysis. The sites were selected in order to contrast the regional vegetation history revealed from the swamp sediments with the local history of the Tsuga forest reconstructed from the Hollow sediments. Specific objective were (1) to document the natural and anthropogenic disturbance history, (2) to examine the long-term vegetation dynamics of the two sites resulting from environmental change, species migration, and disturbance, especially with respect to Tsuga, and (3) to contrast the pre- and post-settlement vegetation and environments. The Swamp and Hollow cores contain continuous sediment records covering the past 12 300 and 9500 yr, respectively. Regional vegetation changes are delimited in six pollen zones: I, Herb zone (12 500-11 800 yr BP); II, Picea zone (11 800-9500 yr BP); III, Pinus-Quercus zone (9350-8350 yr BP); IV, Tsuga-northern hardwoods zone (8350-1750 yr BP); V, Tsuga-Castanea-hardwoods zone (1750-200 yr BP); and, VI Post-settlement zone (200 yr BP-present). No disturbances are detected in the periods of tundra or boreal vegetation from 12 500 to 8 350 yr BP. Since 8350 yr BP three distinct disturbance processes are detectable: (1) fires recorded in discrete charcoal horizons, (2) the apparent pathogenic decline of Tsuga (4700-3500 yr BP) and the blight of Castanea (@? 1915 A.D.), and (3) post-settlement forest cutting, burning, land clearance, and cultivation (1750 A.D. to present). Log-ratio diagrams facilitate the differentiation of local vegetation history in the Tsuga forest from the regional upland history. Tsuga has been the dominant taxon at the Hollow for the last 8000 yr although its abundance has fluctuated with disturbance. Tsuga, Ulmus, Populus, Acer rubrum, A. saccharum, Betula, and Castanea apparently are more abundant locally than in the regional vegetation, which has a larger component of Pinus, Quercus, and Carya. Post-disturbance vegetation dynamics in the Tsuga forest are controlled by the type and intensity of disturbance and the pool of available species, determined by species migration and climate change. Tsuga decreases and subsequently recovers in 300-1200 yr following most disturbances. Northern hardwood taxa (Acer

saccharum, Fagus, Betula spp., Ulmus, Fraxinus) generally decrease after fire and human activity and recover with Tsuga. For the period 8000-3000 yr BP Pinus, Quercus, and occasionally Populus and Nyssa increase following fire and the Tsuga decline at 4700 yr BP. However, since its immigration at @?3000 yr BP Castanea has been the major species to increase upon disturbance of the Tsuga-northern hardwood forest. The study documents the long-term dominance of a forest stand by Tsuga canadensis and northern hardwood species despite repeated infrequent disturbances. The ability of these taxa to gradually reassume their former abundance following disturbance is presumably a consequence of their shade tolerance and longevity and the low frequency of disturbance.

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Oxygen isotope variations spanning the last glacial cycle and the Holocene derived from ice-core records for six sites in Greenland (Camp Century, Dye-3, GRIP, GISP2, Renland and NorthGRIP) show strong similarities. This suggests that the dominant influence on oxygen isotope variations reflected in the ice-sheet records was regional climatic change. Differences in detail between the records probably reflect the effects of basal deformation in the ice as well as geographical gradients in atmospheric isotope ratios. Palaeotemperature estimates have been obtained from the records using three approaches: (i) inferences based on the measured relationship between mean annual

δ18O of snow and of mean annual surface temperature over Greenland; (ii) modelled inversion of the borehole temperature profile constrained either by the dated isotopic profile, or (iii) by using Monte Carlo simulation techniques. The third of these approaches was adopted to reconstruct Holocene temperature variations for the Dye 3 and GRIP temperature profiles, which yields remarkably compatible results. A new record of Holocene isotope variations obtained from the NorthGRIP ice-core matches the GRIP short-term isotope record, and also shows similar long-term trends to the Dye-3 and GRIP inverted temperature data. The NorthGRIP isotope record reflects: (i) A generally stronger isotopic signal than is found in the GRIP record; (ii) several short-lived temperature fluctuations during the first 1500 yr of the Holocene; (iii) a marked cold event at ca. 8.2 ka (the '8.2 ka event'); (iv) optimum temperatures for the Holocene between ca. 8.6 and 4.3 ka, a signal that is 0.6‰ stronger than for the GRIP profile; (v) a clear signal for the Little Ice Age; and (vi) a clear signal of climate warming during the last century. These data suggest that the NorthGRIP stable isotope record responded in a sensitive manner to temperature fluctuations during the Holocene.

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We present oxygen and carbon isotope records of benthic foraminifera from the glacial stage 6 through interglacial substage 5e (Eemian) sections of several cores from the subpolar North Atlantic. The cores range in water depth from 1451 to 2658 m. In one core, we generated estimates of sea surface temperature (SST) and of ice-rafted detritus (IRD) delivery for all of stage 5. We reconstruct bathymetric profiles of delta(13)C for stage 6, Termination II, and three time slices of substage 5e. The delta(13)C profiles indicate that local deep water geometry during stage 6 was similar to that of the last glaciation, with glacial North Atlantic Intermediate Water (GNAIW) overlying deeper waters partially of southern origin. An anomalously large peak in IRD, coupled with planktonic delta(18)O evidence for iceberg melting immediately precedes Termination II and is otherwise similar to stage 2 and 3 Heinrich events. During the termination, low delta(13)C values are observed in cores near and above 2 km, providing evidence of reduced GNAIW production in association with deglacial melting. In the shallowest cores, low delta(13)C values persist into early

substage 5e, indicating that the southward retreat of southern-source waters was not completed until well after the beginning of the substage. In contrast to Termination I, SSTs remained cold until the end of the deglaciation; this may be why there is little evidence from marine and terrestrial sequences for a pronounced climate reversal on Termination II despite what is now clear evidence of a significant reduction in ocean ventilation. The faunal data suggest that SSTs during early 5e were about 7 degrees C warmer than during the glaciation. SST rose several degrees from early to middle substage 5e, peaked in middle substage 5e at about 10 degrees C above glacial values, and then gradually declined by about 5 degrees C. These changes were linked to variations in water mass configuration, as interpreted from benthic delta(13)C evidence. Most of the evidence suggests that oscillations superimposed on the gradual SST trend were 1 degrees-2 degrees C, in contrast to the larger SST changes (3 degrees-4 degrees C) documented for substage 5d.

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Accurate reconstructions of late-Quaternary land-cover change are needed to better understand past interactions of the terrestrial biosphere with other components of the earth system. This paper presents a sequence of reconstructed needleleaved and broadleaved tree-cover densities for North America since the last glacial maximum, generated from fossil-pollen data and present-day tree-cover estimates derived from the Advanced Very High Resolution Radiometer (AVHRR). For this study, a refined form of the modern analog technique was developed, called the hierarchical analog technique, which can constrain paleoenvironmental properties even for fossil-pollen assemblages without close analogs in the modern-pollen record. Pollen taxa from samples that are

compositionally unlike any modern-pollen samples are regrouped into plant functional categories based upon phenology, life form, leaf shape, and climatic tolerances, and the analog analysis rerun. Reclassifying individual pollen taxa into broader functional categories enables analogs to be found when no compositional analogs exist, but at a cost of increased uncertainties in the analog estimates. Tests of the hierarchical analog technique shows that it accurately reconstructs present-day tree-cover densities. The median standard deviation for each individual estimate is 2 concentrations. By 14 ka, broadleaved tree-cover densities had begun to rise in the southeastern US and needleleaved forests grew in the western US, southeastern US, and as a belt along the southern margin of the Laurentide Ice Sheet. By the mid-Holocene, the northern and western needleleaved forests had joined. Needleleaved and broadleaved tree densities continued to increase until European settlement. Mapping percent tree-cover represents a useful alternative to biome-based classification schemes, enabling a fuller representation of vegetational gradients in space and time, and can be directly compared to the tree distributions simulated by dynamic global vegetation models. In effect, by calibrating the modern-pollen data against the AVHRR-derived estimates of tree-cover, the fossil-pollen data are applied to extrapolate satellite-based observations into the Quaternary, enabling study of vegetation dynamics and land-cover change at timescales beyond the period of direct observation by engineered remote sensors. Analog-based approaches, however, require extensive networks of surface and fossil-pollen samples, making further data collection a priority in sparsely sampled regions of the world.

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